

Time series of microphysical structure of a thundercloud examined with hydrometeor classification method for X-band polarimetric radar



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Introduction

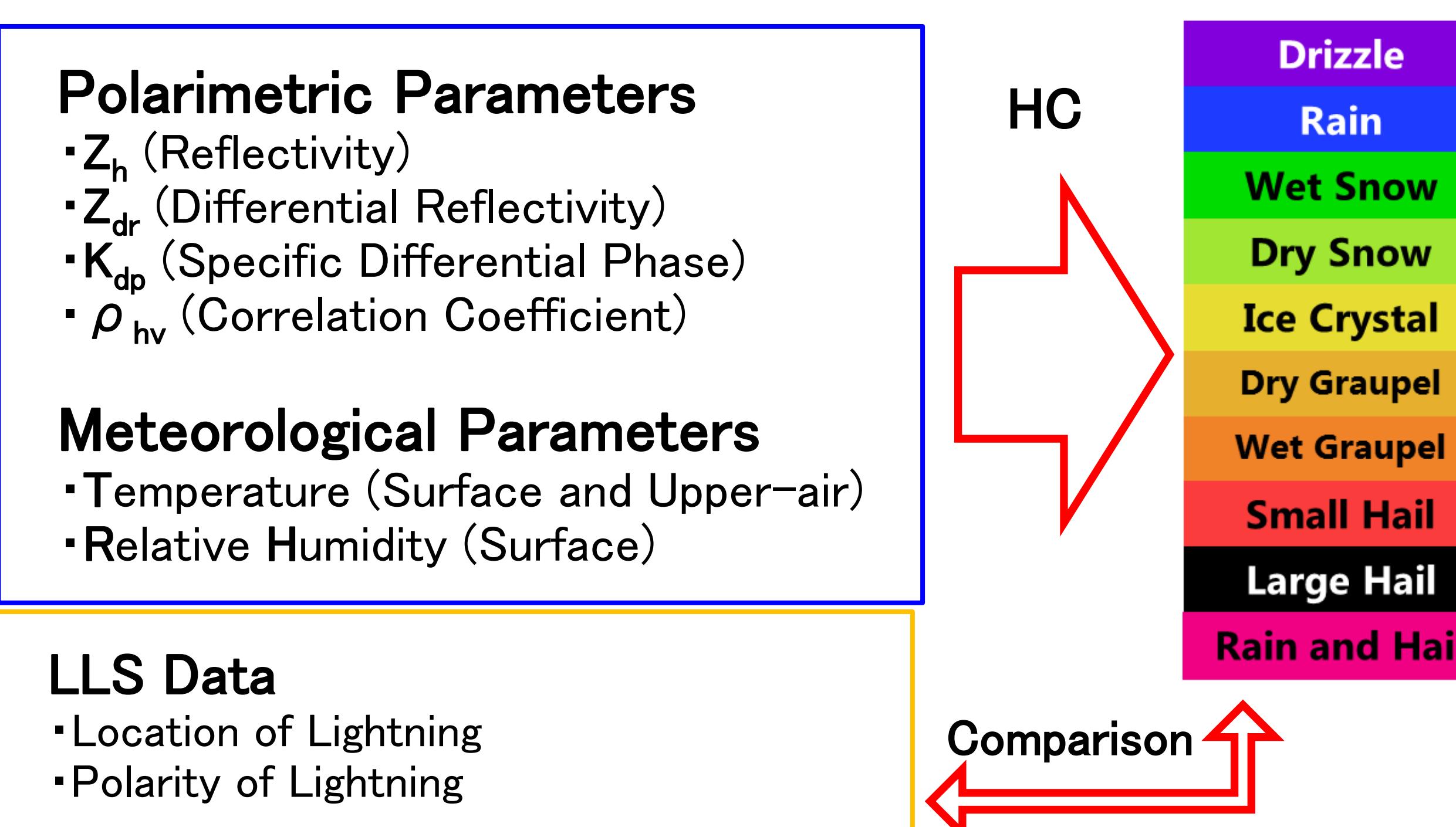
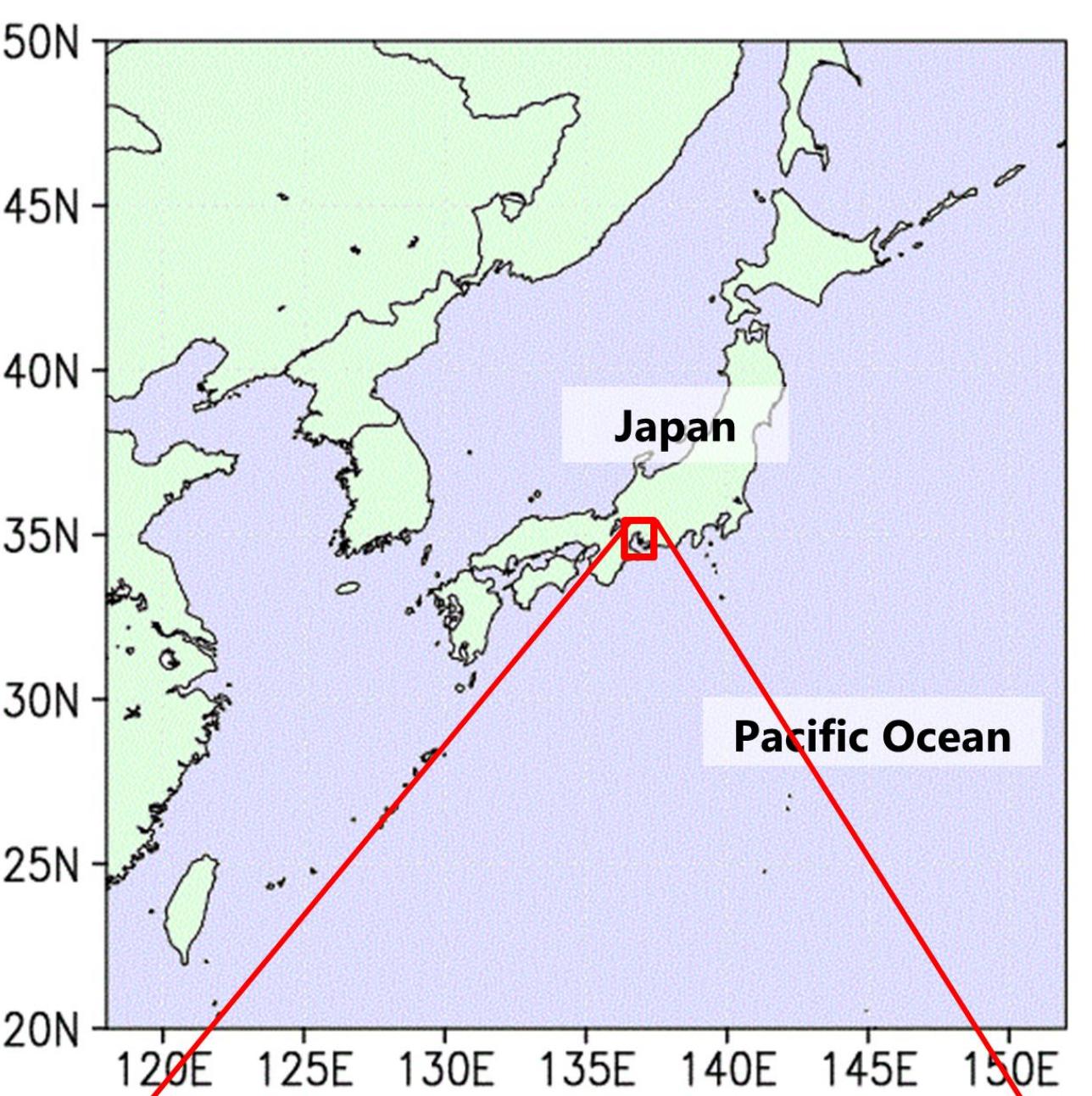
Polarimetric radars are useful instrument to obtain microphysical information and we have modified HC method for S-band polarimetric radar (S-pol) described in Liu and Chandrasekar (2000) to adapt to X-band polarimetric radars (X-pol) and tried hydrometeor classification (hereafter, HC) with X-pol. (Kouketsu and Uyeda, ERAD2010).

To evaluate the HC method, thunderclouds are useful target because solid hydrometeors (snowflake, ice crystal and graupel) are included and their relative locations in the cloud are closely related to the polarity of lightning. In this study, we targeted a single thundercloud of which we could observe entire life cycle in relation to lightning polarity.

Conclusions

- We conducted HC for a single thundercloud with the HC method tuned for X-pol and examined the microphysical structure of the cloud.
- The relation between the volume of graupel (ice crystal) region and the frequency of negative (positive) CG is consistent with the polarity of CG expected from the riming electrification process and, therefore, our HC method can be considered to be reasonable for the single thundercloud.

Data

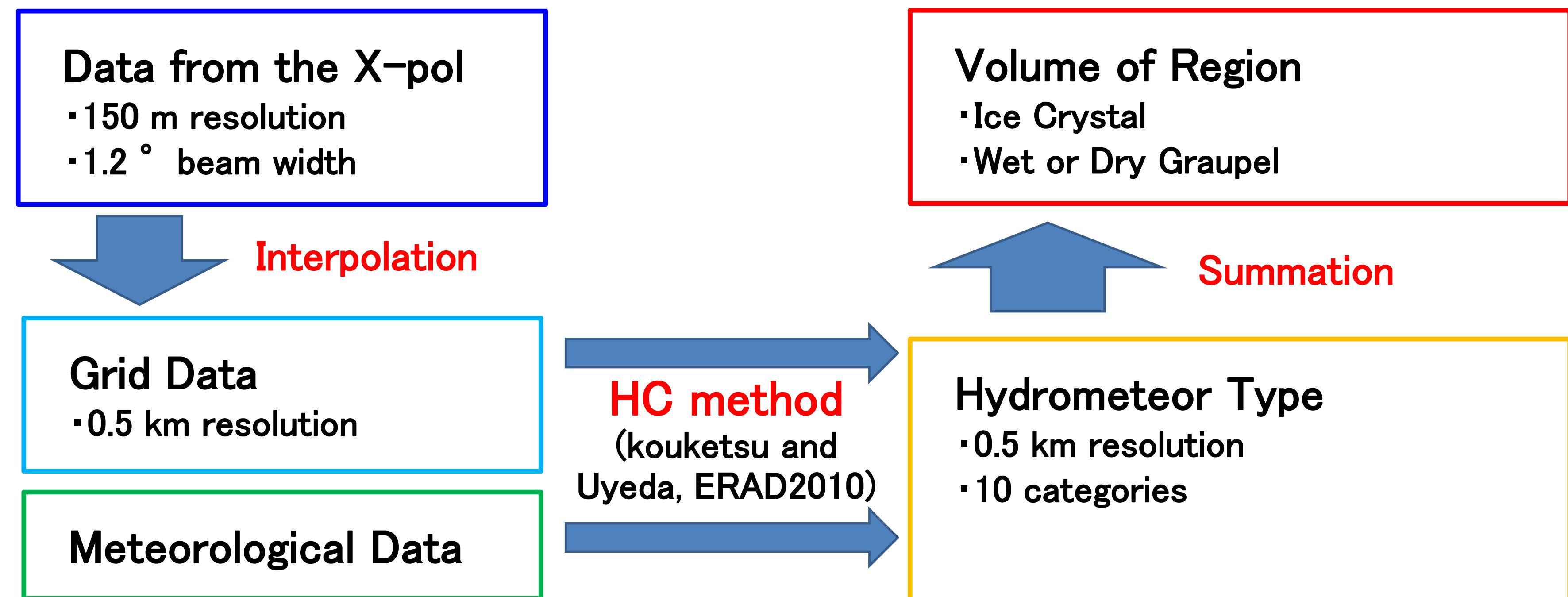


Characteristics of the X-pol

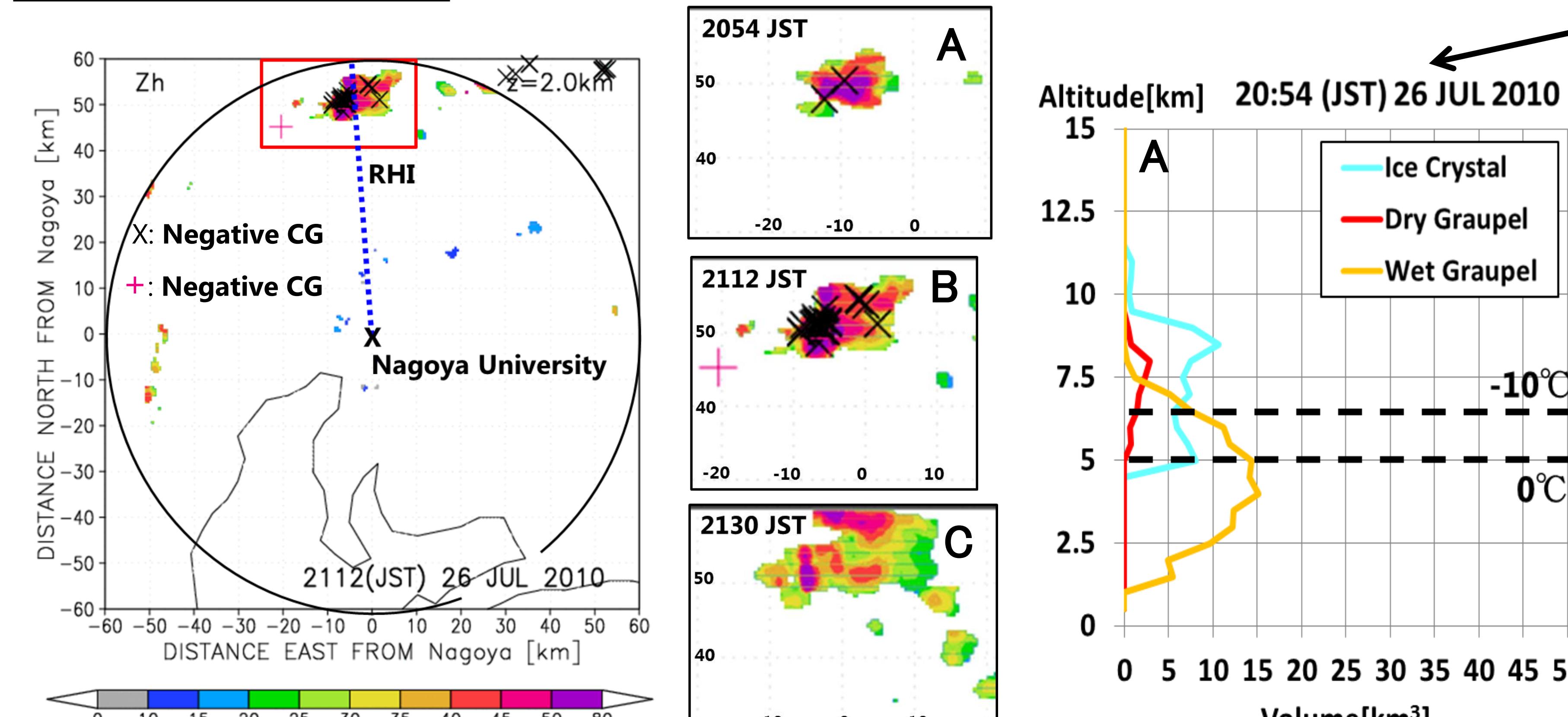
Frequency	9375 MHz
Antenna size	2.0 m
Beam width	1.2°
Transmitter Type	Solid state component
Peak Power	200 W
Max range	61.8 km
Pulse width	1 µs (within 5 km) 32 µs (beyond 5 km, pulse compression)
PRF	2000 Hz / 1600 Hz (dual PRF)
Transmission	45° or H only or V only
Rotation rate	3.0 rpm (PPI), 1.2 rpm (RHI)
Resolution	150 m
Nyquist velocity	16.0 ms ⁻¹ / 12.8 ms ⁻¹

▲ The location of the X-pol

Method

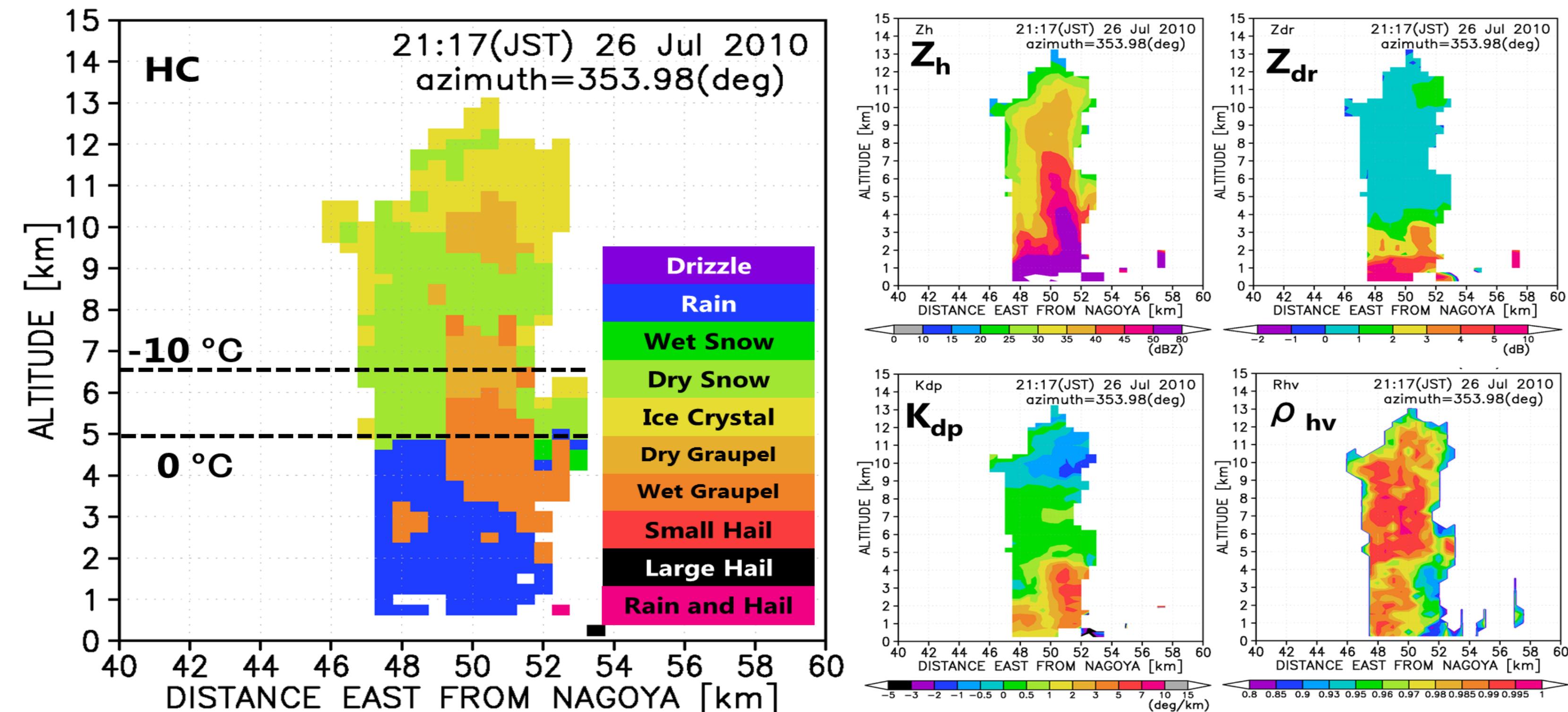


Case Overview



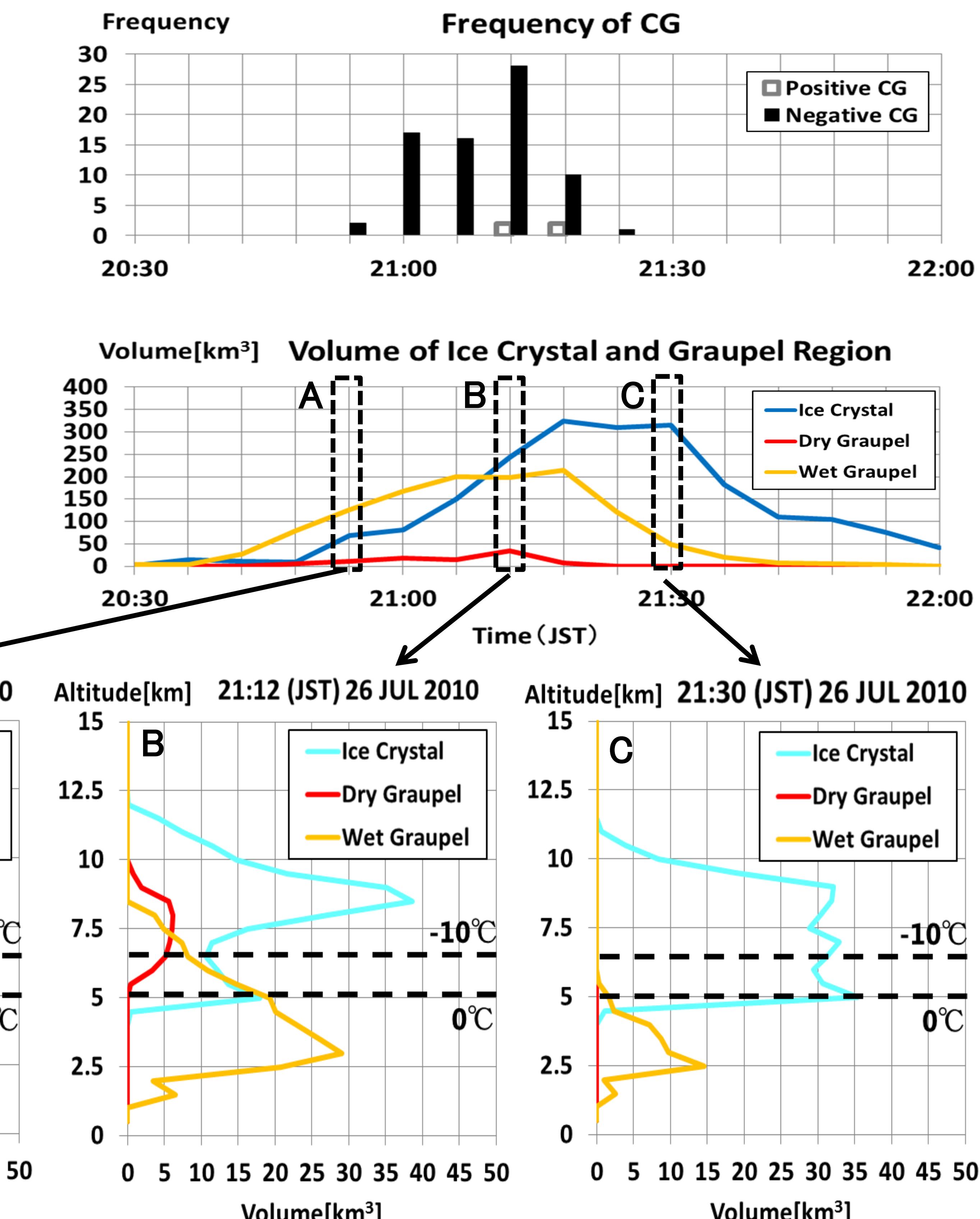
▲ Z_h at 2-km height at 2112 JST on 26 July 2010

▼ The result of HC and polarimetric parameters of RHI at 2117 JST



Characteristics of the X-pol ▲

Result and discussion



▲ Vertical distribution of volume of hydrometeors ▲

- Negative CGs were observed only when there was large volume of wet graupel region (120 km^3).
- There was main dry graupel region around the height where the temperature was -10°C (6.5 km height) when many negative CGs were observed.
- When the volume of ice crystal region increased rapidly and reached the peak, positive CGs were observed.
- These facts are consistent with the polarity of CG expected from the riming electrification process (Takahashi, 1978).